

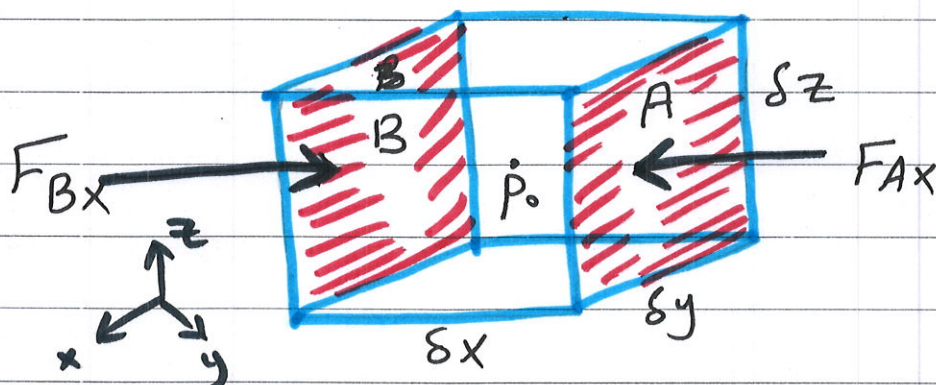
$$\frac{d\vec{V}}{dt} = \underbrace{\rho \vec{G}}_{\text{force}} + \underbrace{\text{Gravitational}}_{\text{force}} + \underbrace{\text{Viscous}}_{\text{"friction"}} - 2\vec{\omega} \times \vec{V} + \vec{\omega} \times (\vec{\omega} \times \vec{r})$$

The Forces in the Equation of motion:

1/ The pressure gradient force.

The pressure force is due to the variation of the atmospheric pressure from point to point. The pressure at a given point and a given time is independent of direction.

Consider the air parcel like the box in the fig. the pressure at the center of box is p_0 .



The volume of the box is $\delta V = \delta x \delta y \delta z$

The pressure at the surface A is

$$p_{Ax} = p_0 + \frac{\partial p}{\partial x} \left(\frac{\delta x}{2} \right).$$

The Force at the surface A is:

$$\vec{F}_{Ax} = - \left[p_0 + \frac{\partial p}{\partial x} \left(\frac{\delta x}{2} \right) \right] \delta y \delta z i$$

The pressure at the surface B is:

$$p_{Bx} = p_0 - \frac{\partial p}{\partial x} \left(\frac{\delta x}{2} \right)$$

The force at the surface B is

$$\vec{F}_{Bx} = \left[p_0 - \frac{\partial p}{\partial x} \left(\frac{\delta x}{2} \right) \right] \delta y \delta z i$$

the net pressure force in the x-direction is

$$\vec{F}_x = \vec{F}_{Bx} - \vec{F}_{Ax}$$

$$\vec{F}_x = \left[p_0 - \frac{\partial p}{\partial x} \left(\frac{\delta x}{2} \right) \right] \delta y \delta z i - \left[p_0 + \frac{\partial p}{\partial x} \left(\frac{\delta x}{2} \right) \right] \delta y \delta z i$$

$$\vec{F}_x = -2 \frac{\partial p}{\partial x} \left(\frac{\delta x}{2} \right) \delta y \delta z i$$

$$\vec{F}_x = - \frac{\partial p}{\partial x} \delta v i$$

$$\delta v = \delta x \delta y \delta z$$

Since $\delta V = \frac{m}{\rho}$ where m is mass
~~and~~ and ρ is density.

$$\vec{F}_x = -\frac{\partial P}{\partial x} \delta V \mathbf{i} = -\frac{m}{\rho} \frac{\partial P}{\partial x} \mathbf{i} \quad \div m$$

$$\vec{f}_x = \frac{\vec{F}_x}{m} = -\frac{1}{\rho} \frac{\partial P}{\partial x} \mathbf{i} \quad x\text{-direction.}$$

Similarly for y and z directions

$$\vec{f}_y = -\frac{1}{\rho} \frac{\partial P}{\partial y} \mathbf{j} \quad y\text{-direction}$$

$$\vec{f}_z = -\frac{1}{\rho} \frac{\partial P}{\partial z} \mathbf{k} \quad z\text{-direction.}$$

The net total of force per unit mass :

$$\vec{f} = \vec{f}_x + \vec{f}_y + \vec{f}_z$$

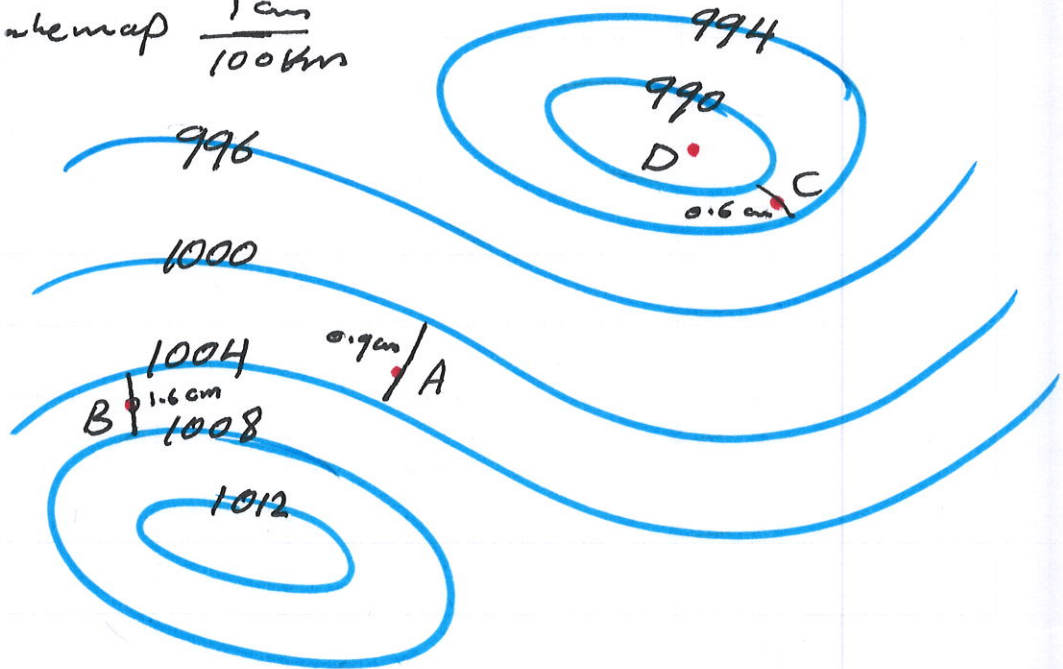
$$\vec{f} = -\frac{1}{\rho} \left[\frac{\partial P}{\partial x} \mathbf{i} + \frac{\partial P}{\partial y} \mathbf{j} + \frac{\partial P}{\partial z} \mathbf{k} \right].$$

$$\vec{f} = -\frac{1}{\rho} \vec{\nabla} P = -\frac{1}{\rho} \cdot \text{grad } P.$$

The direction of the pressure gradient force $\vec{f}$ is in the opposite direction of the pressure gradient $\vec{\nabla} P$, and $\nabla P \approx \frac{\Delta P}{\Delta n}$ where ΔP is the contour interval for the isobars and Δn is the horizontal distance between the isobars.

At the four points shown in the picture below, estimate the magnitude of the acceleration due to the pressure gradient force. Assume a density of 1.23 kg/m^3 . The isobars are labeled in mb.

Scale map $\frac{1 \text{ cm}}{100 \text{ km}}$



$$\vec{f} = -\frac{1}{\rho} \nabla p$$

$$|\vec{f}| = \frac{1}{\rho} \nabla p = \frac{1}{\rho} \frac{\Delta p}{\Delta n}$$

point B:

$$\Delta p = 1008 - 1004 = 4 \text{ mb} = 4 \text{ hPa}$$

$$= 4 \times 10^2 \text{ Pa} = 4 \times 10^2 \text{ N/m}^2$$

$$\Delta n = 1.6 \text{ cm} \times \frac{100 \text{ km}}{1 \text{ cm}} = 160 \text{ km}$$

$$\Delta n = 160 \times 10^3 \text{ m}$$

$$|\vec{f}| = \frac{1}{1.23} \frac{4 \times 10^2}{160 \times 10^3} = 0.002 \text{ m/sec}^2$$

الضغط ينخفض
في اتجاه اليمين

point A :

$$|\vec{f}| = \frac{1}{\rho} \frac{\Delta p}{\Delta n}$$

$$\Delta p = 1004 - 1000 = 4 \text{ mb} = 4 \times 10^2 \text{ N/m}^2$$

$$\Delta n = 0.9 \text{ cm} \times \frac{100 \text{ km}}{1 \text{ cm}} = 90 \text{ km}$$

$$|\vec{f}| = \frac{1}{1.23} \times \frac{4 \times 10^2}{90 \times 10^3} = 0.0036 \text{ m/sec}^2$$

point C :

$$\Delta p = 994 - 990 = 4 \text{ mb} = 4 \times 10^2 \text{ N/m}^2$$

$$\Delta n = 0.6 \text{ cm} \times \frac{100 \text{ km}}{1} = 60 \text{ km}$$

$$|\vec{f}| = \frac{1}{1.23} \times \frac{4 \times 10^2}{60 \times 10^3} = 0.0054 \text{ m/sec}^2$$

point D

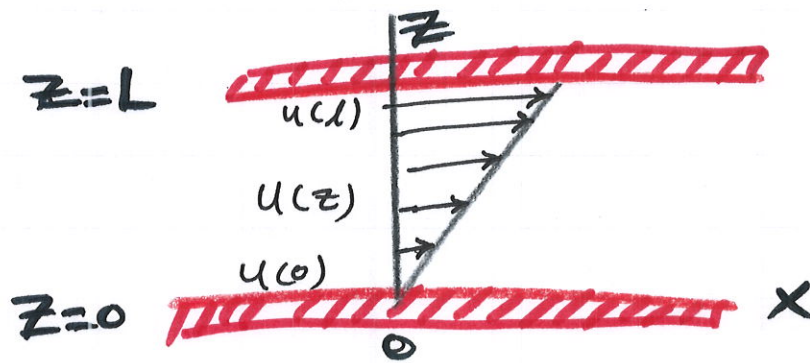
$$|\vec{f}| = \text{Zero.}$$

because $\Delta p = \text{Zero.}$

2- The Viscous Force .

viscous force is due to friction caused by interactions of molecules of a fluid.

If we consider a layer of incompressible fluid between two horizontal plates separated by a distance (L) as shown in the fig.



The layer in contact with the upper and lower plates will move as the following:

at $z=L$ the fluid moves at speed $u(L)=u_0$ and at $z=0$ the fluid is motionless $u(0)=0$.

$$F \propto \frac{A u_0}{L}$$

where u_0 is peak speed, A is area and L is depth.

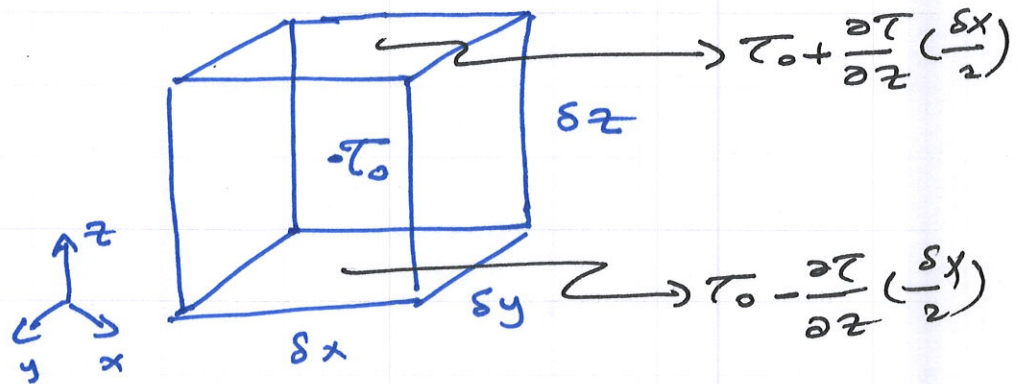
$$\therefore F = \mu \frac{A u_0}{L}$$

where μ dynamical viscosity coefficient.

The shearing stress $\tau = \frac{F}{A}$

$$\text{Thus } \tau_x = \frac{F_x}{A} = \mu \frac{U_0}{L} = \mu \frac{\partial u}{\partial z}$$

From the fig. along the top of the cube the force in x -direction is: $\tau_0 + \frac{\partial \tau}{\partial z} \left(\frac{\delta x}{2} \right)$



on the bottom face of the cube, the force in the x -direction is: $\tau_0 - \frac{\partial \tau}{\partial z} \left(\frac{\delta x}{2} \right)$

$$\text{The net force is: } \vec{F}_x = \left[\sum \vec{\tau}_x \right] A$$

$$\vec{F}_x = \left\{ \left[\tau_0 - \frac{\partial \tau}{\partial z} \left(\frac{\delta x}{2} \right) \right] - \left[\tau_0 + \frac{\partial \tau}{\partial z} \left(\frac{\delta x}{2} \right) \right] \right\} \delta y \delta z \hat{i}$$

$$\vec{F}_x = - \frac{\partial \tau}{\partial z} \delta x \delta y \delta z$$

$$\vec{F}_x = -\frac{\partial \tau}{\partial z} \delta v \quad i$$

~~xxxxxxxxxxxx~~

$$\delta v = \frac{m}{\rho} \Rightarrow \vec{F}_x = -\frac{\partial \tau}{\partial z} \frac{m}{\rho} \quad i$$

$$\therefore \frac{\vec{F}_x}{m} = \vec{f}_x = -\frac{1}{\rho} \frac{\partial \tau}{\partial z} \quad i$$

$$\vec{f}_x = -\frac{1}{\rho} \frac{\partial}{\partial z} \left(\mu \frac{\partial u}{\partial z} \right) \quad i \quad \text{x-direction}$$

$$\text{H.W} \quad \vec{f}_y = -\frac{1}{\rho} \frac{\partial}{\partial z} \left(\mu \frac{\partial v}{\partial z} \right) \quad j \quad \text{y-direction.}$$

$$\text{H.W} \quad \vec{f}_z = -\frac{1}{\rho} \frac{\partial}{\partial z} \left(\mu \frac{\partial w}{\partial z} \right) \quad k \quad \text{z-direction}$$

The total force per unit mass: $\Rightarrow \vec{f} = \vec{f}_x + \vec{f}_y + \vec{f}_z$

$$\vec{f} = -\frac{1}{\rho} \left[\frac{\partial}{\partial z} \left(\mu \frac{\partial u}{\partial z} \right) i + \frac{\partial}{\partial z} \left(\mu \frac{\partial v}{\partial z} \right) j + \frac{\partial}{\partial z} \left(\mu \frac{\partial w}{\partial z} \right) k \right]$$

$$\vec{f} = -\frac{\mu}{\rho} \left[\frac{\partial}{\partial z} \left(\frac{\partial u}{\partial z} \right) i + \frac{\partial}{\partial z} \left(\frac{\partial v}{\partial z} \right) j + \frac{\partial}{\partial z} \left(\frac{\partial w}{\partial z} \right) k \right].$$

$$\vec{f} = -\frac{\mu}{\rho} \frac{\partial}{\partial z} \left[\left(\frac{\partial u}{\partial z} \right) i + \left(\frac{\partial v}{\partial z} \right) j + \left(\frac{\partial w}{\partial z} \right) k \right].$$

$$\vec{f} = -\frac{\mu}{\rho} \frac{\partial}{\partial z} \left[\frac{\partial \vec{V}}{\partial z} \right]. \quad \text{where } \vec{V} = u\vec{i} + v\vec{j} + w\vec{k}.$$

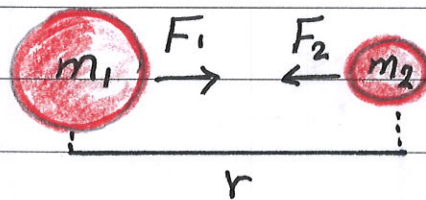
The acceleration due to the viscous force can be written as:

$$\boxed{\vec{f} = -\frac{1}{\rho} \frac{\partial \tau}{\partial z}}$$

The viscosity of the atmosphere is small, so we can ignore it. 20

3.1 The Gravitational force.

Newton's Law of universal gravitation states that "any two element of mass in the universe attracts each other with a force proportional to their masses and inversely proportional to the square of the distance between them".



Newton's Law can be written as a vector equation:

$$\vec{g} = -G \frac{m_1 m_2}{r^3} \vec{r}$$

where $\vec{g}$ attraction of m_1 on m_2 (force of gravitation)

$\vec{r}$ is vector from m_1 to m_2

G is universal gravitational constant $= 6.66 \times 10^{-11} \text{ Nm}^2 \text{ kg}^{-2}$

IF we assume $m_2 = 1 \text{ kg}$

$$\therefore \vec{g} = -G \frac{m_1}{r^3} \vec{r}$$

If $m_1 = M$ 'total mass of earth is equal to 5.988×10^{24} Kg.

The acceleration due to the gravitational force at the surface of earth ($r = a = 6378$ Km) is

$$\vec{g}_* = -G \frac{M}{a^2} \vec{r}$$

At some altitude z above the surface of the earth, the acceleration due to the gravitational force is:

$$\vec{g}_* = -G \frac{M}{(a+z)^2} \vec{r}$$

* $\vec{r}$ is the vector from the center of earth to particle in the atmosphere.

* $\vec{g}_*$ is directed toward the center of earth.

4/ Centrifugal force.

The earth is no inertial system, but a rotating system; therefore Centrifugal and Coriolis force occur.

The term $\omega^2 \vec{R}$ represents the centrifugal acceleration due to the earth's rotation, the centrifugal force is always directed away from the axis of rotation. The centrifugal force is combined with the gravitational force to define a new force called GRAVITY.

The acceleration due to gravity force is defined as :

$$\vec{g} = \vec{g}_* + \omega^2 \vec{R}$$

When we see gravity in an equation of motion, keep in mind that it is a combination of the gravitational acceleration plus centrifugal acceleration.

The gravity in an equation of motion has a plus sign, but keep in mind that if written in component form it lies solely in the negative k-direction.

$$\vec{g} = -g\mathbf{k}.$$

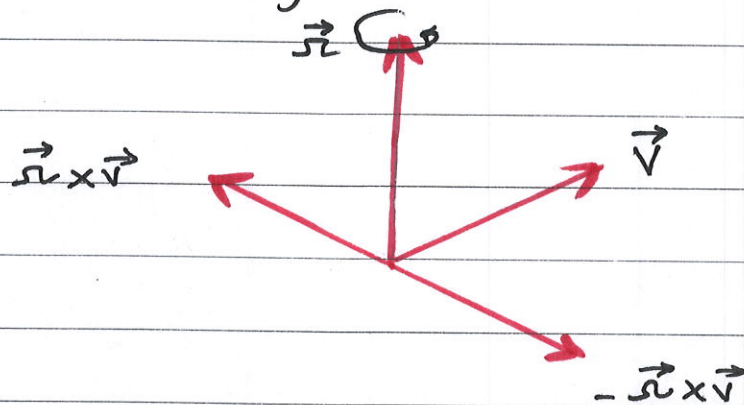
5/ The Coriolis force.

Coriolis force on an air parcel with velocity in a coordinate system with angular velocity $\vec{\omega}$ is:

$$\text{Coriolis force} = -2\vec{\omega} \times \vec{v}$$

The minus sign is important, just like in advection.

The Coriolis force acts perpendicular to the direction of velocity vector:

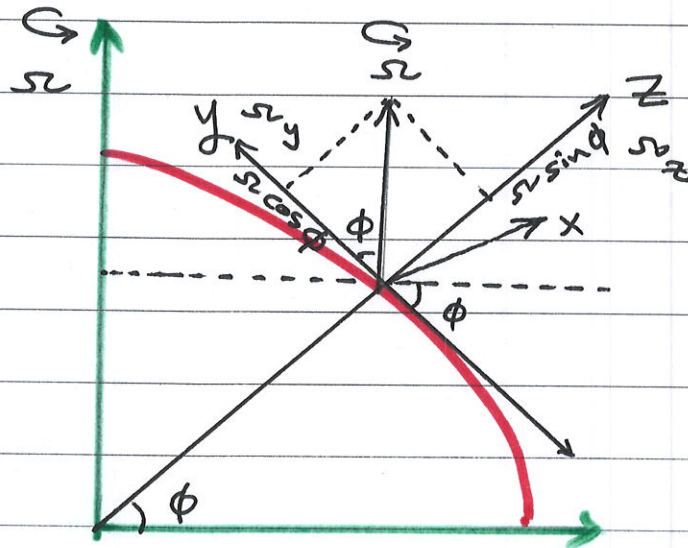


in the northern hemisphere, where ω is counterclockwise, the Coriolis force acts to the right of the velocity vector, it tends to deflect air parcel to the right of their direction of motion.

The Coriolis force changes only the direction of particle's motion, not its speed, because the force is always perpendicular to the direction of motion.

The components of Coriolis force.

From the Fig.



$$\Omega_x = 0$$

$$\Omega_y = \Omega \cos \phi$$

$$\Omega_z = \Omega \sin \phi, \text{ So } \dots$$

$$\text{Coriolis force} = -2 \vec{\Omega} \times \vec{V} = \begin{vmatrix} \mathbf{i} & \mathbf{j} & \mathbf{k} \\ 0 & \Omega_y & \Omega_z \\ u & v & w \end{vmatrix}$$

$$-2 \vec{\Omega} \times \vec{V} = -2(\Omega_y w - \Omega_z v) \mathbf{i} + (0 - \Omega_z u) \mathbf{j} - (0 - \Omega_y u) \mathbf{k}$$

$$-2 \vec{\Omega} \times \vec{V} = (2\Omega_z v - 2\Omega_y w) \mathbf{i} - \Omega_z u \mathbf{j} + \Omega_y u \mathbf{k}$$

$$-2 \vec{\Omega} \times \vec{V} = (2\Omega \sin \phi v - 2\Omega \cos \phi w) \mathbf{i} - \Omega \sin \phi u \mathbf{j} + \Omega \cos \phi u \mathbf{k}$$

$$\text{let } f = 2\Omega \sin \phi \text{ and } e = 2\Omega \cos \phi$$

$$\text{Thus } -2 \vec{\Omega} \times \vec{V} = (fv - ew) \mathbf{i} - fu \mathbf{j} + eu \mathbf{k}$$

The components of the equation of motion:

The Equation of motion is:

$$\frac{d\vec{v}}{dt} = -\frac{1}{\rho} \vec{\nabla} p + \vec{g}_* + \vec{f} - 2\vec{\Omega} \times \vec{v} + \vec{\Omega}^2 \vec{R}$$

We note that $\vec{g}_*$ and $\vec{\Omega}^2 \vec{R}$ are only forces which depend solely on position.

$$\vec{g} = \vec{g}_* + \vec{\Omega}^2 \vec{R}$$

where $\vec{g}$ is gravity.

$$\frac{d\vec{v}}{dt} = -\frac{1}{\rho} \vec{\nabla} p + \vec{g} + \vec{f} - 2\vec{\Omega} \times \vec{v} \dots \dots \textcircled{1}$$

Now we find the x, y and z component of each term in equation ①.

$$\frac{d\vec{v}}{dt} = \frac{du}{dt} \vec{i} + \frac{dv}{dt} \vec{j} + \frac{dw}{dt} \vec{k} \dots \dots \textcircled{A}$$

$$-\frac{1}{\rho} \vec{\nabla} p = -\frac{1}{\rho} \frac{\partial p}{\partial x} \vec{i} - \frac{1}{\rho} \frac{\partial p}{\partial y} \vec{j} - \frac{1}{\rho} \frac{\partial p}{\partial z} \vec{k} \dots \dots \textcircled{B}$$

$$\vec{g} = -g \vec{k} \dots \dots \textcircled{C}$$

$$\vec{f} = -\frac{1}{\rho} \frac{\partial}{\partial z} (\mu \frac{\partial u}{\partial z}) \vec{i} - \frac{1}{\rho} \frac{\partial}{\partial z} (\mu \frac{\partial v}{\partial z}) \vec{j} - \frac{1}{\rho} \frac{\partial}{\partial z} (\mu \frac{\partial w}{\partial z}) \vec{k} \dots \dots \textcircled{D}$$

$$-2\vec{\Omega} \times \vec{v} = (fv - ew) \vec{i} - fu \vec{j} + ev \vec{k} \dots \dots \textcircled{E}$$

put equations A, B, C, D and E in Equation (1).

$$\begin{aligned} \frac{du}{dt}i + \frac{dv}{dt}j + \frac{dw}{dt}k = & -\frac{1}{\rho} \frac{\partial p}{\partial x}i - \frac{1}{\rho} \frac{\partial p}{\partial y}j - \frac{1}{\rho} \frac{\partial p}{\partial z}k \\ & -gk + \left(-\frac{1}{\rho} \frac{\partial}{\partial z} (\mu \frac{\partial u}{\partial z})\right)i - \frac{1}{\rho} \frac{\partial}{\partial z} (\mu \frac{\partial v}{\partial z})j - \frac{1}{\rho} \left(\frac{\partial}{\partial z} (\mu \frac{\partial w}{\partial z})\right)k \\ & + (fv - ew)i - fu j + eu k. \end{aligned}$$

The components of this equation are:

$$\frac{du}{dt} = -\frac{1}{\rho} \frac{\partial p}{\partial x} - \frac{1}{\rho} \frac{\partial}{\partial z} (\mu \frac{\partial u}{\partial z}) + fv - ew \quad x\text{-Dir.}$$

$$\frac{dv}{dt} = -\frac{1}{\rho} \frac{\partial p}{\partial y} - \frac{1}{\rho} \frac{\partial}{\partial z} (\mu \frac{\partial v}{\partial z}) - fu \quad y\text{-Dir.}$$

$$\frac{dw}{dt} = -\frac{1}{\rho} \frac{\partial p}{\partial z} - \frac{1}{\rho} \frac{\partial}{\partial z} (\mu \frac{\partial w}{\partial z}) - g + eu \quad z\text{-Dir.}$$

To assess which terms can be neglected, we assign an "order of magnitude" to all the variables and parameters in the equations.

For synoptic scales the following order of magnitude are: -

name	symbol	order of magnitude
Horizontal velocity	U	10 m/sec.
vertical velocity	W	0-01 m/sec.
Horizontal distance	L	1000 km = 10^6 m
vertical distance	H	10 km = 10^4 m
Horizontal pressure	Δp	10 mb
density	ρ	1 kg/m^3
Time	$\frac{L}{U}$	1 day (10^5 sec)
viscosity	μ	$1.46 \times 10^{-5} \frac{\text{m}^2}{\text{sec}}$
omega	ω	$7-3 \times 10^{-5} \text{ sec}^{-1}$
latitude	ϕ	45°
radius of earth	R	$6.378 \times 10^6 \text{ m}$.

Using these scales and parameters, in the x-momentum equation we have the following order of magnitude.

$$\frac{du}{dt} = -\frac{1}{\rho} \frac{\partial p}{\partial x} - \frac{1}{\rho} \frac{\partial}{\partial z} \left(\mu \frac{\partial u}{\partial z} \right) + f v - e w.$$

$$\frac{\partial u}{\partial t} = \frac{U^2}{L} = \frac{10^2}{10^6} = 10^{-4} \text{ m/sec}^2$$

$$\frac{1}{\rho} \frac{\partial p}{\partial x} = \frac{\Delta p}{\rho L} = \frac{10^3}{1 \times 10^6} = 10^{-3} \text{ m/sec}^2$$

$$\frac{1}{\rho} \frac{\partial}{\partial z} \left(\mu \frac{\partial u}{\partial z} \right) = \mu \frac{\partial^2 u}{\partial z^2} = \frac{\nu \frac{\partial^2 u}{\partial z^2}}{H^2} = \frac{1.5 \times 10^{-5} \times 10}{10^8}$$

$$\therefore \frac{1}{\rho} \frac{\partial}{\partial z} \left(\mu \frac{\partial u}{\partial z} \right) = 10^{-12} \text{ m/sec}^2$$

$$f v = 2 \Omega v \sin \phi = 2 \Omega U \sin \phi = 2 \times 7.3 \times 10^{-5} \times 10 \times \sin 45^\circ$$

$$f v = 10^{-3} \text{ m/sec}^2$$

$$e w = 2 \Omega W \cos \phi = 2 \times 7.3 \times 10^{-5} \times 10^{-2} \times 0.7$$

$$e w = 10^{-6} \text{ m/sec}^2.$$

A similar analysis for y and z equations.

Many of terms are very small compared to others, and can therefore be ignored without much loss of accuracy. We can therefore ignore the viscous terms and Coriolis term that involves the vertical velocity.

Thus the components of equation of motion are:

$$\frac{du}{dt} = -\frac{1}{\rho} \frac{\partial p}{\partial x} + f v$$

$$\frac{dv}{dt} = -\frac{1}{\rho} \frac{\partial p}{\partial y} - f u$$

$$\frac{\partial p}{\partial z} = -\rho g \rightarrow \text{Hydrostatic Equation.}$$

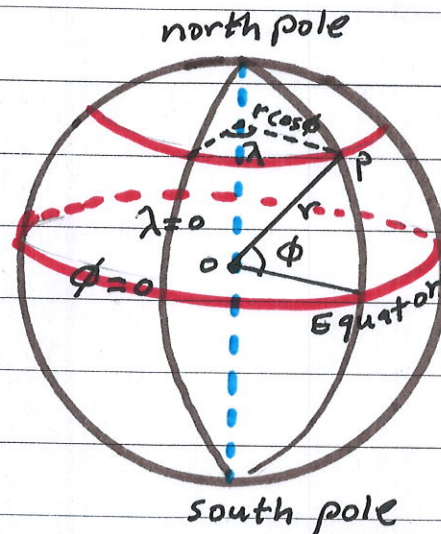
Hydrostatic equation is expressing a balance between gravity acting downwards and the vertical component of the pressure force acting upwards.

This system of equation is quite often used in analysis of dynamical problems.

Coordinate system (spherical system).

It is partial to use a spherical coordinate system with origin at the center of the earth to describe dynamical aspects.

The coordinate system is rotating together with the planet.



λ longitude in easterly direction.

ϕ latitude in northerly direction.

z geometric altitude.

R radius of planet $r = R + z$.

$u = r \cos \phi \frac{d\lambda}{dt}$ zonal component of wind $\vec{V}$
measured towards the east.

$v = r \frac{d\phi}{dt}$ meridional component of wind $\vec{V}$
measured towards the north.

$w = \frac{dz}{dt}$ vertical component of $\vec{V}$.

Equation of motion in spherical coordinates.

The whole set in spherical coordinates without friction.

$$\frac{du}{dt} = \frac{uv}{r} \tan \phi - \frac{uw}{r} + 2\Omega \sin \phi v - 2\Omega \cos \phi w - \frac{1}{r \cos \phi} \frac{\partial p}{\partial \lambda}$$

$$\frac{dv}{dt} = -\frac{u^2}{r} \tan \phi - \frac{vw}{r} - 2\Omega \sin \phi u - \frac{1}{r} \frac{\partial p}{\partial \phi}$$

$$\frac{dw}{dt} = \frac{u^2 + v^2}{r} + 2\Omega \cos \phi u - g - \frac{1}{r} \frac{\partial p}{\partial r}$$

In the Lagrangian frame we have to take care of the total derivative.

$$\frac{d}{dt} = \frac{\partial}{\partial t} + \frac{u}{r \cos \phi} \frac{\partial}{\partial \lambda} + \frac{v}{r} \frac{\partial}{\partial \phi} + w \frac{\partial}{\partial r}$$