

Example 1:

The recession limb of a flood hydrograph is given below. The time is indicated from the arrival of peak. Assuming the interflow component to be negligible, estimate the base flow and surface flow recession coefficients. Also, estimate the storage at the end of day-3.

Time from peak (day)	Discharge (m <sup>3</sup> /s)	Time from peak (day)	Discharge (m <sup>3</sup> /s)
0	90	4.0	3.8
0.5	66	4.5	3.0
1.0	34	5.0	2.6
1.5	20	5.5	2.2
2.0	13	6.0	1.8
2.5	9.0	6.5	1.6
3.0	6.7	7.0	1.5
3.5	5		

Solution:

The data are plotted on semi-log paper. The flow from 4.5 to 7 days are seen to lie on straight line AB in Fig(9).??

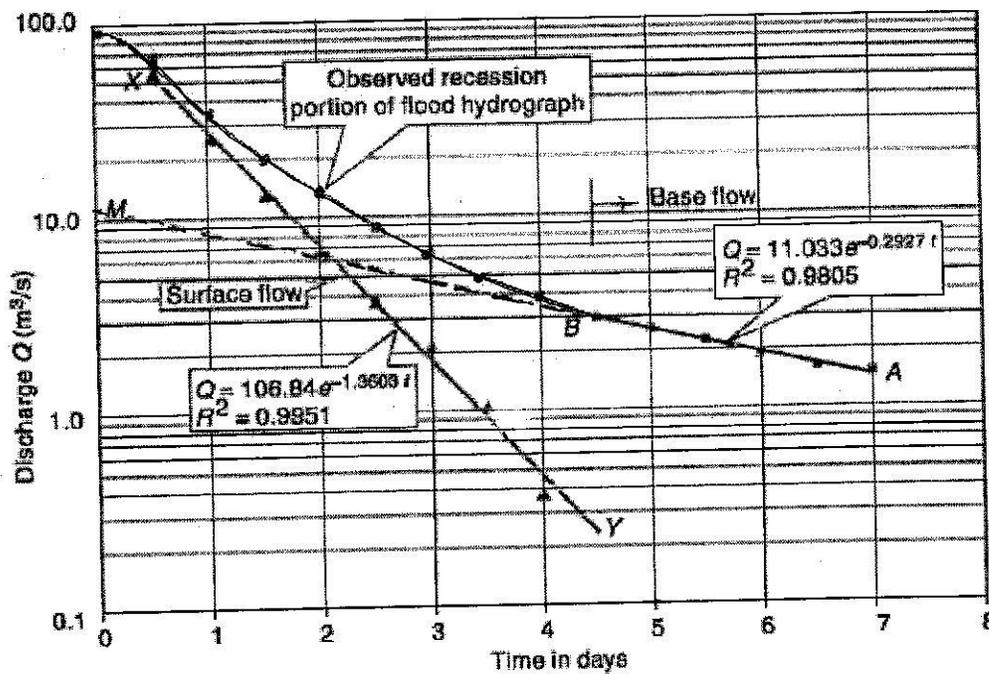


Fig. 9 Storage Recession Curve - Example 1

From eqn:  $\frac{\phi_2}{\phi_1} = Kr^{(t_2 - t_1)}$

$\log Krb = \frac{1}{t_2 - t_1} \log \left( \frac{\phi_2}{\phi_1} \right)$  for base flow (Line BM).

Select: at  $t_1 = 0 \rightarrow \phi_b = 11$   
 at  $t_2 = 4.5 \rightarrow \phi_b = 3$  }  $\Rightarrow \log Krb = \frac{1}{4.5 - 0} \log \left( \frac{3}{11} \right)$   
 $Krb = 0.749 \left[ \phi_b = 11 (0.749)^t \right]_{BM}$

Thus; the surface runoff depletion is obtained by subtracting the base flow from the given recession limb of the hydrograph. and apply the above procedure taking at  $t_1 = 0, \phi_s = 79$  and

$t_2 = 3 \text{ day} \rightarrow \phi_s = 2.078 \text{ m}^3/\text{s} \Rightarrow Ks = 0.297$   
 $\Rightarrow \left[ \phi_s = 79 (0.297)^t \right]_{XY}$

From eq (6): The storage available at end of  $t$ -days is:

$$S_b + S_s)_t = \left( \frac{Q_b}{-\ln K_b} + \frac{Q_s}{-\ln K_s} \right)$$

at period  $t = 3$  days:

$$Q_b = 11 (0.749)^3 = 4.622 \text{ m}^3/\text{s}$$

$$Q_s = 79 (0.297)^3 = 2.07 \text{ m}^3/\text{s}$$

$$\Rightarrow S_3 = \frac{4.622}{-\ln 0.749} + \frac{2.07}{-\ln 0.297} = 16 + 1.7 = 17.7 \frac{\text{m}^3}{\text{s}} \cdot \text{day (cumec-day)}$$

$$= 1.52928 \text{ Mm}^3$$

هذا الحجم من المياه المتوفرة أثناء من نصف اليوم السابق

H.W In a stream the base flow is observed to be  $30 \text{ m}^3/\text{s}$  on May 1<sup>st</sup> and  $23 \text{ m}^3/\text{s}$  on May 10<sup>th</sup>. If there is no rain during May, estimate the base flow on 30 May and the volume of ground water storage on May 1<sup>st</sup> & May 30<sup>th</sup>.

Effective rainfall (ER)

الطرالوثر

Effective rainfall or Excess rainfall (ER) is the part of rainfall that becomes direct runoff at the out let of catchment area.

It is thus, the total rainfall from which abstractions (infiltration & initial loss) are subtracted. The DRH was correlated with so-called hyetograph of effective rainfall (ERH) shown in Fig (10). DRH & ERH represent same quantity but in different units.

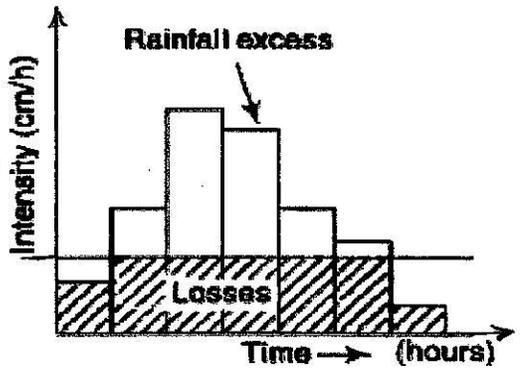


Fig. 10 Effective Rainfall Hyetograph (ERH)

$\phi$ -index, represents the average rainfall intensity above which the rainfall volume is equal to the runoff volume.

Example 2: A storm produced a direct runoff of  $5.8 \text{ cm}$ . The data of the storm is given below. Estimate the  $\phi$ -index.

Time (hr)	0	2	4	6	8	10	12	14	16
Acc. depth (cm)	0	0.4	1.3	2.8	5.1	6.9	8.5	9.5	10

Solution:

Time (hr)	0 - 2	2 - 4	4 - 6	6 - 8	8 - 10	10 - 12	12 - 14	14 - 16
depth (cm)	0.4	0.9	1.5	2.3	1.8	1.6	1	0.5
Intensity (cm/hr)	0.2	0.45	0.75	1.15	0.9	0.8	0.5	0.25

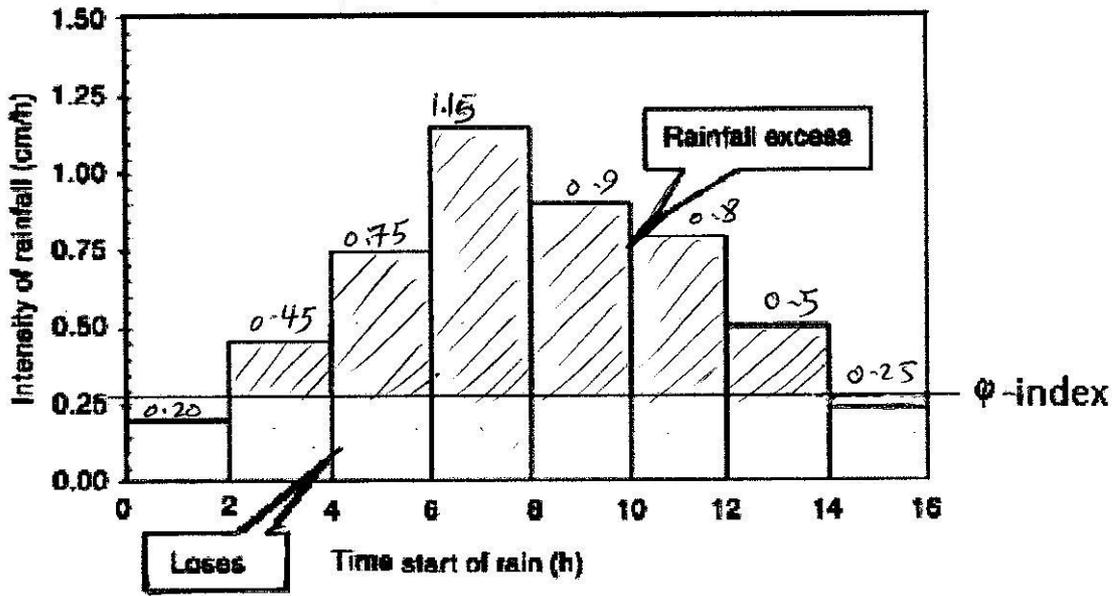


Fig. 11 Hyetograph and Rainfall Excess of the Storm - Example 2

① assume  $0.45 < \phi < 0.5$  ( $M=5$ )

$\sum (I_i - \phi) \Delta t = R_d$        $\Rightarrow R_d$ : direct Runoff

$\sum (I_i \Delta t - \phi \Delta t) = \sum_{i=1}^5 I_i \Delta t - 5 \phi \Delta t = 5.8$   
 $= 0.75(2) + 1.15(2) + 0.9(2) + 0.8(2) + 0.5(2) - 5\phi(2) = 5.8$   
 $\Rightarrow \phi = 2.4 \text{ cm/hr}$  (Assuming is not correct)

② assume  $0.2 < \phi < 0.25$  ( $M=7$ )

$\Rightarrow 5.8 = 0.45(2) + 0.75(2) + 1.15(2) + 0.9(2) + 0.8(2) + 0.5(2)$   
 $+ 0.25(2) - 7\phi(2)$   
 $\Rightarrow \phi = 0.271 > 0.25$  (Not ok)

③ assume  $0.25 < \phi < 0.45$  ( $M=6$ )

$\Rightarrow 5.8 = 0.45(2) + 0.75(2) + 1.15(2) + 0.9(2) + 0.8(2) + 0.5(2) - 6\phi(2)$   
 $\Rightarrow \phi = 0.275$  (OK.  $0.25 < \phi < 0.45$ )

لاحظ ان الخوارزمية قائمة على فرضية ان نسبة اولية المطر غير قاي  $\phi$   
 يتم التحقق من النسبة الفرضية عبر تطبيق مفهوم المطر المباشر  
 هو كائناً لغت السبع المباشر

Example 3

Rainfall of magnitude 3.8 cm and 2.8 cm occurring on two consecutive 4-h durations on a catchment of area 27 km<sup>2</sup> produced the following hydrograph of flow at outlet of the catchment. Estimate the rainfall excess and  $\phi$  index.

Time from start of rainfall (h)	-6	0	6	12	18	24	30	36	42	48	54	60	66
Observed flow (m <sup>3</sup> /s)	6	5	13	26	21	16	12	9	7	5	5	4.5	4.5

Solution:

The hydrograph is plotted as in Fig (12). The separation of the base flow of the storm Hydrograph was considered by using simple straight line.

However, DRH starts at t=0, having peak at t=12h and ends at t=48h, i.e., 36h from peak while this period calculated from eq (3) is;

$N = 0.83(27)^{0.2} = 1.6 \text{ days} = 38.5 \text{ h}$   
 (which one is recommended to use?)

It obvious that the base flow has a constant value of 5 m<sup>3</sup>/sec -

The Area under DRH curve is calculated by Trapezoidal scheme

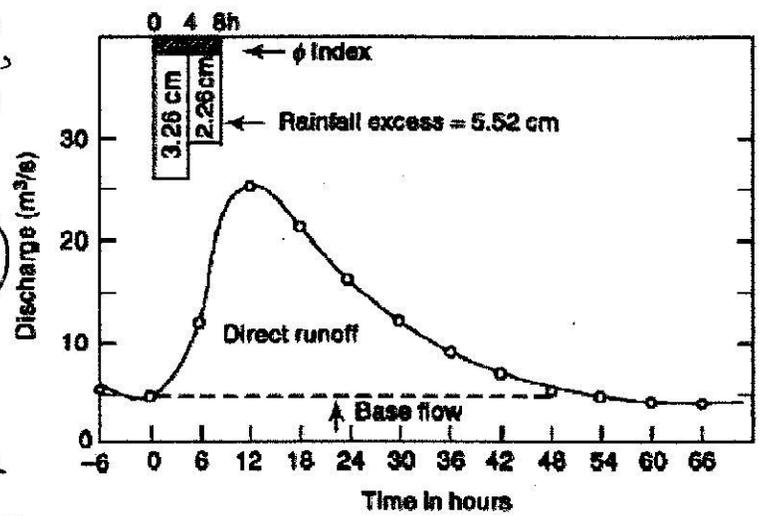


Fig. 12 Base Flow Separation—Example 3

Time (hr)	0	6	12	18	24	30	36	42	48	54	60	66
DRH m <sup>3</sup> /sec	0	8	21	16	11	7	4	2	0			

Area of DRH =  $\Delta t \left[ \frac{0+8}{2} + \frac{8+21}{2} + \frac{21+16}{2} + \frac{16+11}{2} + \frac{11+7}{2} + \frac{7+4}{2} + \frac{4+2}{2} + \frac{2+0}{2} \right]$   
 $= 6 \times 3600 [8+21+16+11+7+4+2] = 1,4904 \times 10^6 \text{ m}^3$

Runoff depth (Rd) =  $\frac{\text{Volume of Runoff}}{\text{Catchment Area}} = \frac{1,4904 \times 10^6}{27 \times 10^6} = 0.0552 \text{ m} = 5.52 \text{ cm excess rainfall}$

Total rainfall = 3.8 + 2.8 = 6.6 cm

$\phi = \frac{6.6 - 5.52}{8} = 0.135 \text{ cm/hr}$

$d_1 = \left[ \frac{3.8}{4} - \phi \right] \times 4 = 3.26 \text{ cm}$  ;  $d_2 = \left[ \frac{2.8}{4} - \phi \right] \times 4 = 2.26 \text{ cm}$  ;  $d_1 + d_2 = Rd$